

Preliminary data on the petrology and K-Ar dating of the Oligocene volcano Briastovo, Eastern Rhodopes

Yotzo Yanev, Zoltan Pecskay

Abstract. The Briastovo volcano, located in Borovitza region of the Eastern Rhodope volcanic area, is over 25 km in diameter and comprises lava flows of a total thickness of over 1500 m. They form 3 volcanic complexes (from bottom to top): phenoandesites (of unknown age and chemistry); medium to coarse porphyric ultrapotassic latites (33.4 ± 1.3 Ma) and quartz-latite domes (32 ± 1.2 Ma). The latter two complexes are part of the II intermediate phase of the Eastern Rhodope volcanism. The latite dome Dolni-Voden, product of the III intermediate phase (31 ± 1.2 Ma), is exposed in the northern periphery of the volcano. Later, the Briastovo volcano was covered by the dome-cluster of Haskovo mineral springs of the III acid phase, composed of three trachyrhyodacite-trachyrhyolite domes with perlitic periphery (29 to 30.4 ± 1.2 Ma).

Key words: ultrapotassic latites, Eastern Rhodopes, Tertiary volcanism.

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Разположеният в Боровишкия район на Източнородопската вулканска област Брястовски вулкан е с диаметър над 25 km и е изграден от лавови потоци с обща дебелина над 1500 m. Тесъставят 3 вулкански комплекса (отдолу нагоре): на феноандезитите (с неизвестна възраст и химизъм), на средно до едропорфирните ултракалиеви латити (възраст 33.4 ± 1.3 млн. г.) и на кварц-латитовите куполи (32 ± 1.2 млн. г.). Последните два комплекса са от II среднокисела фаза на източнородопския вулканизъм. В северната периферия на вулкана се разкрива латитовият купол Долни Воден от III среднокисела фаза (възраст 31 ± 1.2 млн. г.). По-късно върху Брястовския вулкан се разполага вулканският ареал Хасковски минерални бани от III кисела фаза, състоящ се от три трахириодацит-трахириолитови куполи на изстискване с перлитова периферия (възраст от 29 до 30.4 ± 1.2 млн. г.).

Ключови думи: ултракалиеви латити, Източни Родопи, терциерен вулканизъм

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The Briastovo volcano is situated at the Eastern Rhodopes volcanic area, Borovitza region. This is a very large volcano with diameter over 25 km (from the road Novakovo village - Hut "Sini Vrah" in the west to the Susam village in the east

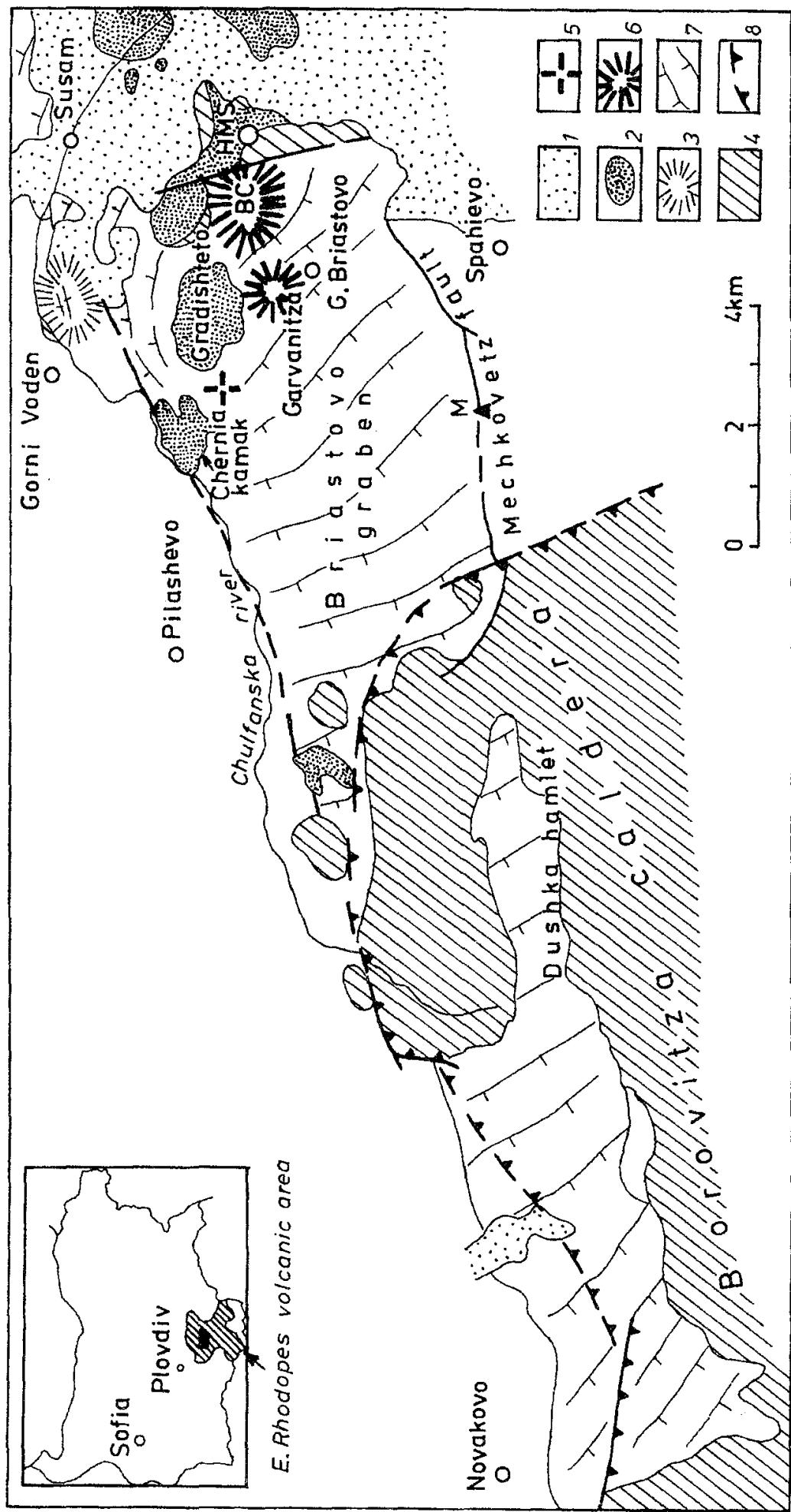


Fig. 1. Schematic geological map of the Briastovo volcano and the NE part of the Borovitza caldera (after Geological Map..., 1989 with additions): 1 — Neogene-Quaternary cover. Oligocene: 2 — acid domes (III acid phase), 3 — latite dome (III intermediate phase), 4 — acid lavas and pyroclastics (II acid phase), 5 — Briastovo intrusion, 6 — quartz-latite domes, 7-ultrapotassic latite lava flows (6 and 7 are of II intermediate phase), 8 — caldera faults. In white - latites not discussed in the text. HMS — Haskovo mineral springs, BC — Banska — chuka dome, M — Mechkovez summit

Table 1

Chemical composition of the magmatic rocks of Briastovo volcano

Таблица 1

Химичен състав на магматитите от Бриастовския вулкан

No Analyses	Intermediate volcanics										Briastovo intrusion				Acid volcanics			
	1 1596b	2 1580	3 1581b	4 1580	5 1670	6 1672	7 1668	8 1668	9 1656	10 1656	11 1651	12 1592g	13 1591g	14 137	15 12j			
SiO ₂	58.46	57.48	58.22	59.34	57.14	58.24	58.22	60.54	59.96	62.83	70.36	70.74	65.08	75.36	75.90			
TiO ₂	0.65	0.68	0.60	0.59	0.66	0.65	0.65	0.59	0.64	0.53	0.38	0.31	0.37	0.10	0.80			
Al ₂ O ₃	14.63	14.43	14.63	13.84	15.01	15.67	14.72	14.87	14.78	13.76	14.30	13.93	13.76	11.54	12.50			
Fe ₂ O ₃	5.46 ^t	3.69	4.62	5.95 ^t	6.21 ^t	2.32	6.41 ^t	6.23 ^t	5.64 ^t	4.71 ^t	2.40	2.24	2.69	1.08	1.08			
FeO	n.d. ^t	2.14	1.27	n.d. ^t	n.d.	2.53	n.d. ^t	n.d. ^t	n.d. ^t	n.d. ^t	0.84	n.d.	n.d.	0.80	0.29			
MnO	n.d.	0.14	0.11	0.12	n.d.	0.16	n.d.	n.d.	n.d.	n.d.	0.03	n.d.	n.d.	n.d.	n.d.			
MgO	4.18	3.87	4.07	3.66	3.25	3.44	3.19	2.67	2.53	2.49	0.53	0.76	0.76	1.33	0.28			
CaO	6.09	5.83	5.18	6.18	5.13	5.76	5.47	3.76	5.09	4.73	1.50	2.38	2.45	1.40	0.52			
Na ₂ O	2.67	2.61	2.01	2.46	2.13	2.45	2.55	2.69	3.80	3.35	3.34	3.51	3.17	2.43	3.80			
K ₂ O	5.25	6.04	6.20	5.61	6.47	5.76	5.20	5.56	5.05	4.78	4.62	5.53	4.47	5.04	4.80			
P ₂ O ₅	n.d.	0.79	0.85	0.66	n.d.	0.65	0.66	0.58	0.55	0.49	n.d.	n.d.	n.d.	0.07	0.04			
CO ₂	n.d.	1.47	n.d.	n.d.	n.d.	0.83	n.d.	n.d.	n.d.	1.17	n.d.	n.d.	n.d.	n.d.	n.d.			
H ₂ O ⁻	n.d.	0.67	n.d.	n.d.	0.35	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	0.82	0.15			
H ₂ O ⁺	2.39	0.07	1.76	1.94	3.94	1.31	2.05	1.54	0.99	0.98	1.50	1.38	6.59	1.08	0.72			
Total	99.78	100.18	99.52	100.35	99.92	99.12	99.03	99.03	99.82	99.80	100.78	99.91	99.82	100.82				

Note: The index "t" indicates all iron given as Fe₂O₃.

Locality: 1 — east of Dushka hamlet, 2 — west of Mechkovets summit (after Radonova, 1973), 3 — the same place, 4 — west of Gradišteto summit; borehole 939, 45 m (slightly altered), 5 — the same borehole, 326 m (moderately altered), 6 — the same place, borehole 970, 557 m (slightly altered), 7 — Yurtlenski-cheshmi (slightly altered); quartz-lattice: 8 — west of Banska-chuka summit; latite: 9 — south of Gorni-Voden village; quartz-montzonite (with anhydrite): 10 — west of Gradišteto summit, borehole 965, 1502 m; trachyhydodacites: 11 — Chernia-kamak summit, 12 — Gradišteto summit, borehole 974, 378 m, 13 — same place, borehole 940, 0 m (perlite); thrahyhydolites: 14 — north and 15 — east of Haskovo mineral springs.

Задележка: Индексът "t" обозначава общо желязо, дадено като Fe₂O₃.

Място: 1 — изт. от с. Душка, 2 — зап. от вр. Мечковец (Радонова, 1973), 3 — същото място, 4 — зап. от вр. Градището, сонд. 939, 45 м (слабо променен), 5 — същия сондаж, 326 м (средно променен), 6 — същото място, сонд. 970, 557 м (слабо променен), 7 — Юртленски чешми (слабо променен); кварцлатит: 8 — зап. от вр. Банска чука; латит: 9 — южно от с. Горни Воден; кварцмонтзонит (с анхидрит); зап. от вр. Градището, сонд. 965, 1502 м; трахириодакти: 11 — Черният камък, 12 — Градището, сонд. 974, 378 м, 13 — същото място, сонд. 974, 378 м, 13 — същото място, сонд. 940, 0 м (перлит); трахириолит: 14 — сев. и 15 — зап. от Хасковски минерални бани.

- Fig. 1). Its western part forms the fundament of the acid volcanic in the north-eastern sector of the Borovitza caldera and its eastern part fills the Briastovo graben with N-NE direction (9×5 km) in the caldera rim.

This is a lava volcano and it consists of three volcanic units in the Briastovo graben (from the bottom to the top): 1) phenoandesite lava flows, 2) coarsely to medium porphyric ultrapotassic latite lava flows, and 3) quartz-latite domes. The first and second units were formed by several hundred metres thick complexes of great number of lava flows, separated by the subaerial epiclastic beds (red siltstone, coals, etc) some metres thick. Based on the borehole data its total thickness is more than 1500 m in the central part of the volcano.

The first unit occurs in the depth of the volcano with a thickness about 500-800 m. Chemical and radiometric data are not available because the volcanics are strongly altered (K u n o v et al., 1997). They contain phenocrysts of plagioclases, clinopyroxene and biotite. In the 1500 m depth they are cut by one quartz-monzonite body — the Briastovo intrusion — constituted by plagioclase ($An_{40-45.5} Or_{2.6-5.1}$), K-feldspar ($Or_{85-86} Ab_{11-14} Cn_{0-1.5}$), biotite ($Mg/Mg+Fe=65-67$ and $Al^{IV}=2.29-2.42$) and actinolitized

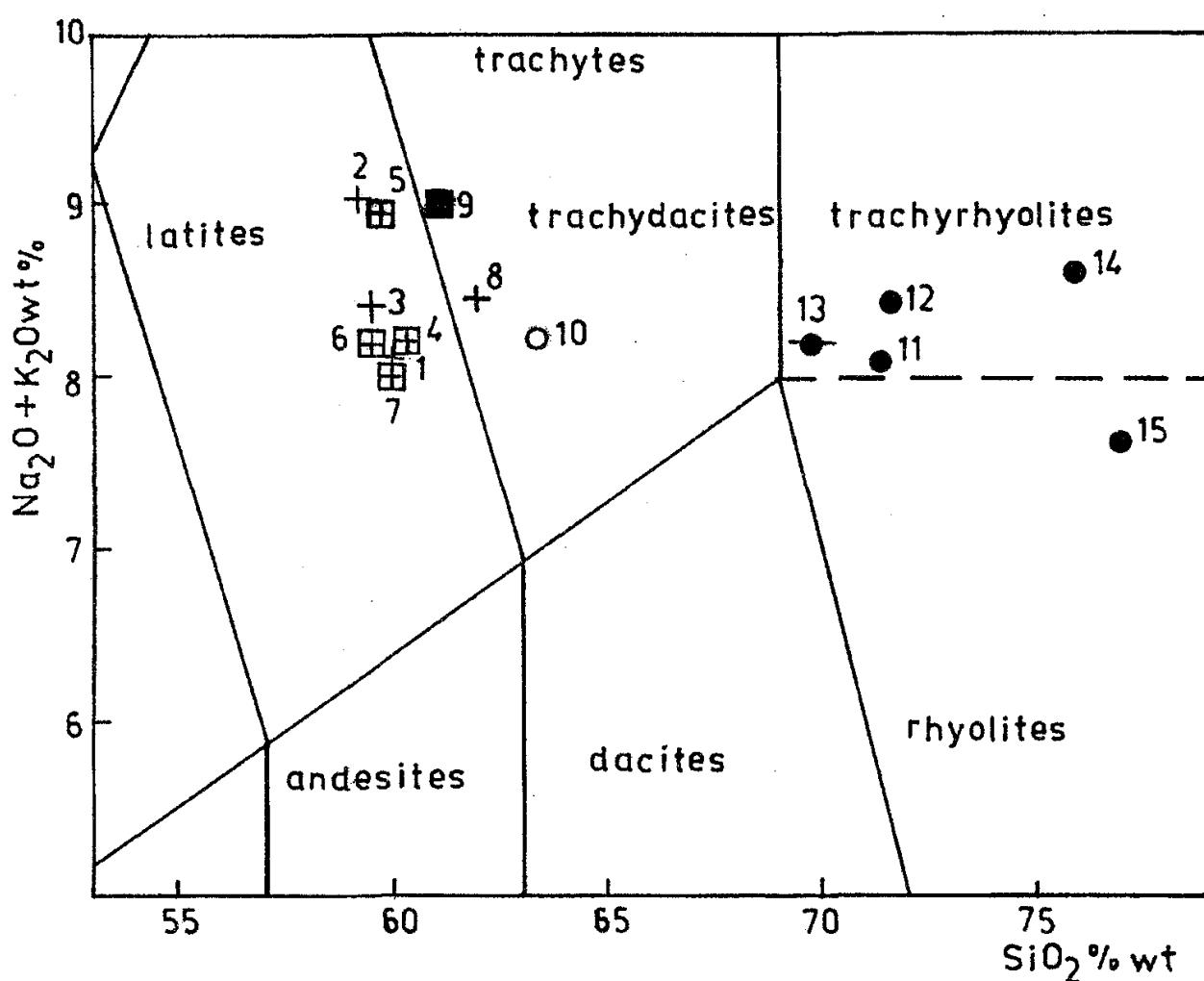


Fig. 2. TAS diagram of IUGS commission (L e M a i t r e, 1989; the limit between rhyolites and trachyrhyolites is according to M y a s h i r o, 1978 and Klassifikatzia..., 1981). For the symbols see Fig. 3; the numbers correspond to Table 1

Фиг. 2. TAS диаграма на IUGS комисия (L e M a i t r e, 1989; границата между риолитите и трахириолитите е според M y a s h i r o, 1978 и Klassifikatzia..., 1981). За знаците на фигуративните точки вж. фиг. 3, а номерата им отговарят на табл. 1

pyroxene. The feldspars crystallisation temperature (according to Fuhrmann, Lindley, 1988, geothermometry) is 635°C.

The second unit is 800 to 1000 m thick and it appears on the surface. The volcanics are fine porphyric (up to 0.2-0.3 cm) in the depth and medium to coarsely porphyric on the surface. In some parts of the flows the sanidine phenocrysts are up to 1-2 cm, rarely up to 5 cm. The volcanics are also altered (Kunov et al., 1997) but preserved in some places, where its composition (Table 1) is of ultrapotassic rocks (according to the Foley et al., 1988, criteria), i.e. with $K_2O > 3\text{wt}\%$, $K_2O/Na_2O > 2$ and $MgO > 3\text{wt}\%$. These rocks were described as trachytes by Hussak in 1883 and by Bontschef in 1896 (mentioned in Ivanov, 1960), but the Na_2O content is very low for a trachyte (Table 1, Fig. 2 and 3). The volcanics contain phenocrysts of sanidine

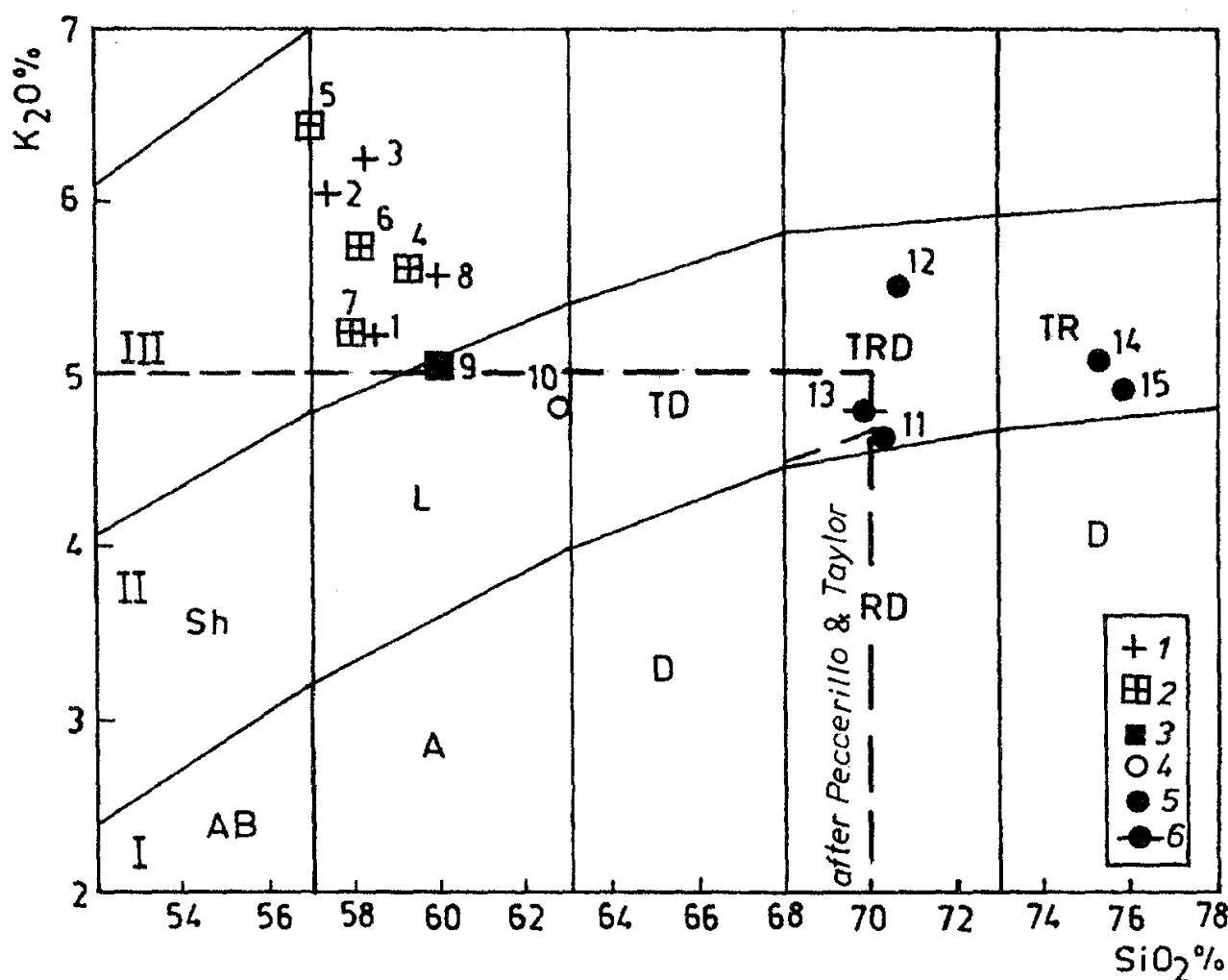


Fig. 3. Peccerillo, Taylor K_2O vs SiO_2 diagram (extended by Stanisheva-Vassileva, Yanev and Harkovska in: Dabovski et al., 1991): I and 2 — ultrapotassic latites (2 — slightly altered), 3 — latite, 4 — quartz-monzonite, 5 — trachyrhyodacites and trachyrhyolites, 6 — perlite (calculated at 100% dry mass). Series: I — High-K Ca-alkaline (BA — basaltic andesite, A — andesite, D — dacite, RD — rhyodacite, R — rhyolite), II — Shoshonitic (Sh — shoshonite, L — latite, TRD — trachyrhyodacite, TR — trachyrhyolite), III — High-K shoshonitic. For the location of the analyses see Table 1

Фиг. 3. K_2O/SiO_2 диаграмма на Peccerillo, Taylor (разширена от Станисева-Василева, Янев и Харковска в: Dabovski et al., 1991): I и 2 — ултракалиеви латити (2 — слабо променени), 3 — латит, 4 — кварцмонционит, 5 — трахириодакити и трахириолити, 6 — перлит (преизчислен към 100% сухо вещество). Серии: I — Висококалиева Ca-алкална (ВА — андезитобазалти, А — андезити, D — дацити, RD — риодакити, R — риолити), II — Шошонитова (Sh-шошонит, L — латити, TD — трахидакити, TRD — трахириодакити, TR — трахириолити), III-висококалиева шошонитова. За мястото на анализите виж табл. 1

($\text{Or}_{62-70}\text{Ab}_{25-28}\text{Cn}_{2-3}$), zoned plagioclase (core $\text{An}_{86-88}\text{Or}_{3-4}$, periphery $\text{An}_{38-43}\text{Or}_{6-7}$), Na diopside-augite ($\text{Wo}_{44-45}\text{En}_{43-44}$ with $\text{Na}_2\text{O}=1.25-1.66\%$ wt), biotite to phlogopite ($\text{Mg}/\text{Mg+Fe}=61-81$, $\text{Al}^{\text{IV}}=2.36-2.47$) in the fine crystalline groundmass with microlites of sanidine with variable composition (from $\text{Or}_{63}\text{Ab}_{30}$ to $\text{Or}_{84}\text{Ab}_9$), anorthoclase ($\text{Or}_{20}\text{Ab}_{70}$), plagioclase and clinopyroxene; accessories: Ti-magnetite, apatite and zircon. The feldspars crystallization temperature (according to Fuhrmann, Lindley, 1988, geothermometry) is 837°C . The age is given in the Table 2.

The volcanic activity was finished by the formation of two quartz-latite domes (Garvanitza and Banska-Chuka ones) some hundred meters in diameter. They contain phenocrysts of sanidine ($\text{Or}_{64-72}\text{Ab}_{2-2.5}\text{Cn}_{1-3}$), zoned plagioclase (core $\text{An}_{52}\text{Or}_3$, periphery $\text{An}_{43-49}\text{Or}_{4-5}$), diopside ($\text{Wo}_{47}\text{En}_{39-40}$), biotite ($\text{Mg}/\text{Mg+Fe}=60-61$, $\text{Al}^{\text{IV}}=2.50-2.56$) and subphenocrysts of quartz. The feldspars crystallization temperature (according to Fuhrmann, Lindley, 1988, geothermometry) is 827°C . The age is given in the Table 2.

At the northern edge of the graben, in the valley of the Chiuflianska River appears the Dolny-Voden latite dome of more recent age. This dome cuts the acid tuffs of the cover of the Briastovo volcano. These latites contain the diopside-augite ($\text{Wo}_{43-44}\text{En}_{43-45}$), plagioclase ($\text{An}_{45-46}\text{Or}_{6-7}$) and phlogopite ($\text{Mg}/\text{Mg+Fe}=77$, $\text{Al}^{\text{IV}}=2.27$) phenocrysts.

Table 2

K-Ar age data for the Briastovo volcanics (analysed at the Institute of Nuclear Research of Hungarian Academy of Sciences - ATOMKI, Debrecen)

Таблица 2

K-Ar данни за възрастта на Брястовските вулканити (анализирани в Института за ядрени изследвания към Унгарската академия на науките - ATOMKI, Дебрецен)

No Analyses	Volcano or dome (locality)	Rock name	K (%)	$^{40}\text{Ar}_{\text{rad}}$ (%)	$^{40}\text{Ar}_{\text{rad}}$ (cc Stp/g)	K/Ar age (Ma)	Volcanic phase (Lower Oligocene)
1670	Briastovo volcano (Yurtlenski-Cheshmi)	ultra-potassic latite	4.88	84.9	6.384×10^{-6}	33.4 ± 1.3	II intermediate
B-24*	Briastovo volcano (1km NE of Mechkovetz summit)	ultra-potassic latite	5.30	—	—	31.5 (average)	II intermediate
1672	Briastovo volcano (Garvanitza dome)	quartz-latite	4.79	89.9	5.997×10^{-6}	31.9 ± 1.2	II intermediate
1668	Dolny-Voden dome	latite	4.78	78.3	5.808×10^{-6}	31.0 ± 1.2	III intermediate
511	Chernia-Kamak dome	trachyrhyo-dacite	4.67	82.2	5.509×10^{-6}	30.2 ± 1.2	III acid
1592g	Gradishteto dome	trachyrhyo-dacite	4.34	86.7	5.162×10^{-6}	30.4 ± 1.1	III acid
12j	Haskovo Mineral Springs dome	trachyrhyo-dacite	4.39	80.5	4.964×10^{-6}	28.8 ± 1.1	III acid

*According to Liov et al. (1989).

During the latter acid phase along the northern border fault of the graben three trachyrhyodacite-rhyolite domes with perlitic periphery (Chernia-kamak, Gradishteto and Haskovo mineral springs) were intruded. They form Haskovo mineral springs dome-cluster (Yanev et al., 1983). The volcanics contain phenocrysts of quartz, sanidine ($\text{Or}_{62-80}\text{Ab}_{20-34}\text{Cn}_{1-4}$), plagioclase ($\text{An}_{18-39}\text{Or}_{3-6}$) and biotite ($\text{Mg/Mg+Fe}=46-65$, $\text{Al}^{IV}=2.22-2.50$); accessories: magnetite, apatite, zircon and titanite. The trachyrhyodacites of Gradishteto dome contain also amphibole (edenite to Mg-hornblende) and diopside-augite ($\text{Wo}_{44-47}\text{En}_{39-48}$). The feldspar crystallization temperature (according to Fuhrmann, Lindsey, 1988, geothermometry) in the Gradishteto dome is 692-699°C and in the Haskovo mineral springs dome is 714-750°C. The age is given in Table 2.

The Briastovo volcanics without being defined as a volcanic edifice are considered of Priabonian age (Lilov et al., 1987) or of Oligocene age - I or II intermediate phase (Ivanov, 1960, 1972; Bojanov, Mavrudchiev, 1961; Yanev, 1989; Geological Map..., 1989). Latter acid domes are referred to the II acid Oligocene phase (Ivanov, 1960, 1972; Yanev et al., 1983; Geological Map..., 1989).

The geochronologic study of the Briastovo volcano is aimed at providing a time framework for related petrological and volcanological studies of volcanic evolution. K-Ar age determinations from whole rock samples are reported from 6 localities in the studied area. Together with the previously obtained data (Lilov et al., 1987) they show a range of 33-29 Ma. The anallitycal data are summarized in Table 2, where their K-Ar age is compared with the age of the Eastern Rhodopes volcanic phases according to Lilov et al. (1987).

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