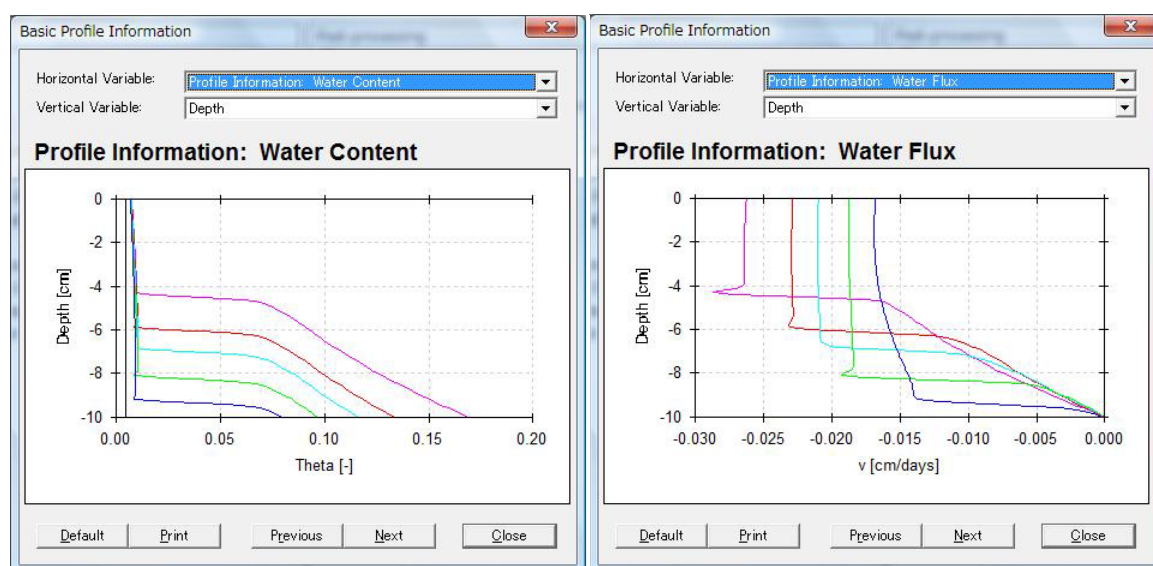


HYDRUS-1D の新機能の紹介と今後の展開



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主催: 農業農村工学会土壌物理研究部会

会場: 三重大学 三翠ホール

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HP1: A COUPLED NUMERICAL CODE FOR VARIABLY SATURATED WATER FLOW, SOLUTE TRANSPORT, AND BIOGEOCHEMISTRY IN SOIL SYSTEMS

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Mie University

Irrigation & Drainage Society, Soil Physics Meeting
October 19, 2008

OUTLINE

- Introduction: Why Coupled Reactive Transport Models?
- HP1: HYDRUS1D - PHREEQC
- Application of the code

What is a Coupled numerical code for reactive transport modeling?

Coupled numerical code ➡ **HP1**

Hydrological code:
water, heat, and
solute transport



HYDRUS

Biogeochemical code:
chemical reactions



PHREEQC



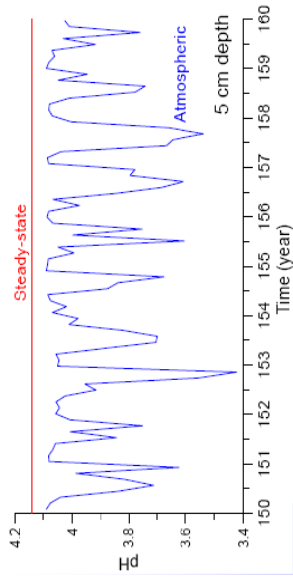
Cycling of major and trace elements in soil-plant systems

- Heterogeneous physical/chemical properties
- Water flow under rainfall – evaporation conditions
- Cation exchange processes
- Surface complexation
- Precipitation/dissolution
- Volatilization
- Root growth and water uptake
- Uptake of metals by plants

Soils are open systems

- Thus: changing geochemical conditions do not only result from static geochemical reactions (“batch”)
- But also: from
 - Amount of available soil water
 - Changing groundwater table
 - Geochemical conditions of water flow

Transient flow conditions $\Rightarrow$ transient geochemical conditions (*example*)

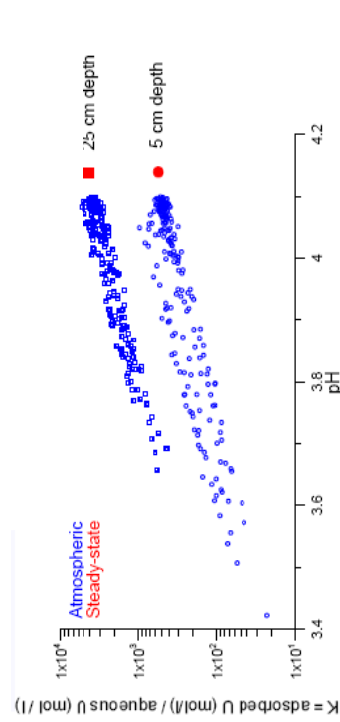


pH variations of soil water under the atmospheric condition:

- water content variations induce pH variations (dry soil $\Rightarrow$ low pH)
- pH variations $\Rightarrow$ variations in sorption potential (low pH $\Rightarrow$ low sorption)

(from HP1 Training Course handouts, 2006, Mol, Belgium)

pH effects on the distribution coefficient K_d for Uranium adsorption (*example*)



(from HP1 Training Course handouts, 2006, Mol, Belgium)

At least one order of magnitude variation in K

Process understanding and assessment of pollutant migration

- Requires a coupled reactive transport model (RTM)
 - Water flow
 - Multicomponent solute transport
 - Biogeochemical processes and reactions (e.g., nitrogen cycle, carbon cycle, etc.)
 - Interaction with biosphere (vegetation), groundwater, surface water (e.g., run off)

What are fully coupled models?

Fully coupled models contain two independent tools:

- Groundwater flow models (e.g. HYDRUS, VS2DI, MODFLOW, etc.)
- and**
- Chemical reaction models (e.g. PHREEQC, MINTEQA2, EQ3NR, etc.)

Why fully coupled models and not geochemical models?

- Adding diffusive and dispersive transport
- Adding physical /spatial/ heterogeneity (e.g. hydraulic parameters)
- Coupling with time-dependent processes as heat transfer, evolving medium properties
- Providing information on the spatial distribution of variables and processes

Why fully coupled models and not classical hydrological models?

- Adding mechanistic models for sorption and their effect on chemical mobility
- Adding effects of aqueous complexation and new species on chemical mobility
- Conversion of contaminants via biochemical reactions in parallel and sequential reaction networks
- Adding dissolution/precipitation and their effect on adsorption

What is HP 1?

HP 1

A Coupled Numerical Code
(Jacques and Šimůnek, 2005)



HYDRUS-1D

(Šimůnek et al., 1998):

Computer program for simulating water, heat and solute movement in one-dimensional variably saturated media.

PHREEQC

(Parkhurst and Appelo, 1999):

Computer program for simulating biogeochemical reactions.

HYDRUS-1D features

- Variable saturated water flow
- Solute transport
- Heat transport
- Root water uptake

PHREEQC features

- Aqueous complexation
- Cation exchange
- Surface complexation – diffuse double layer model and non-electrostatic surface complexation model
- Precipitation/dissolution
- Chemical kinetics
- Redox reactions
- Biological reactions

PHREEQC: Fundamentals

PHREEQC is:

A computer program for biogeochemical reactions, based on the *law of mass action*.

Example: For the reaction : $\text{CaSO}_4 \cdot 2\text{H}_2\text{O} = \text{Ca}^{2+} + \text{SO}_4^{2-} + 2\text{H}_2\text{O}$

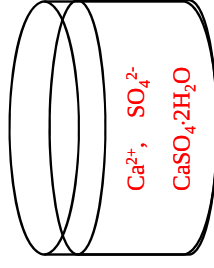
applying *the law of mass action* gives: $K = \frac{[\text{Ca}^{2+}][\text{SO}_4^{2-}][\text{H}_2\text{O}]^2}{[\text{CaSO}_4 \cdot 2\text{H}_2\text{O}]} = 10^{-4.6}$

where K is the equilibrium constant and “[] ” represents concentration

PHREEQC Example: Solubility of Gypsum



adding 1 liter water
(at 25°C)

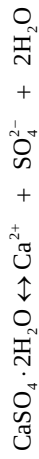


Q1: If the solution is at saturated state, what is the concentration of SO_4^{2-} (or Ca^{2+}) in the liquid phase?

Q2: If 0.001 mol CaCl_2 is added to the saturated solution, what is the SO_4^{2-} concentration in equilibrium?

Q3: If repeat the addition of 0.001mol CaCl_2 10 times, what will be the equilibrium SO_4^{2-} concentration?

Results



Low solubility of gypsum leads to the solubility product as

$$K = \frac{[\text{Ca}^{2+}][\text{SO}_4^{2-}][\text{H}_2\text{O}]^2}{[\text{CaSO}_4 \cdot 2\text{H}_2\text{O}]} = 10^{-4.6} \Rightarrow K = [\text{Ca}^{2+}][\text{SO}_4^{2-}] = 10^{-4.6}$$

Q1: Substituting $[\text{SO}_4^{2-}] = [\text{Ca}^{2+}]$ into the solubility product results in:

$$K = [\text{Ca}^{2+}][\text{SO}_4^{2-}] = [\text{SO}_4^{2-}]^2 = 10^{-4.6} \Rightarrow [\text{SO}_4^{2-}] = 0.00501 \text{ (mol/l)}$$

Results

Q2: Same amount of Δx for Ca^{2+} and SO_4^{2-} precipitates at the new equilibrium. Hence the new concentrations can be given as

$$[\text{Ca}^{2+}] = (0.00501 + 0.001 - \Delta x) \text{ mol/l, and } [\text{SO}_4^{2-}] = (0.00501 - \Delta x) \text{ mol/l}$$

Substituting these concentration into the solubility product leads to

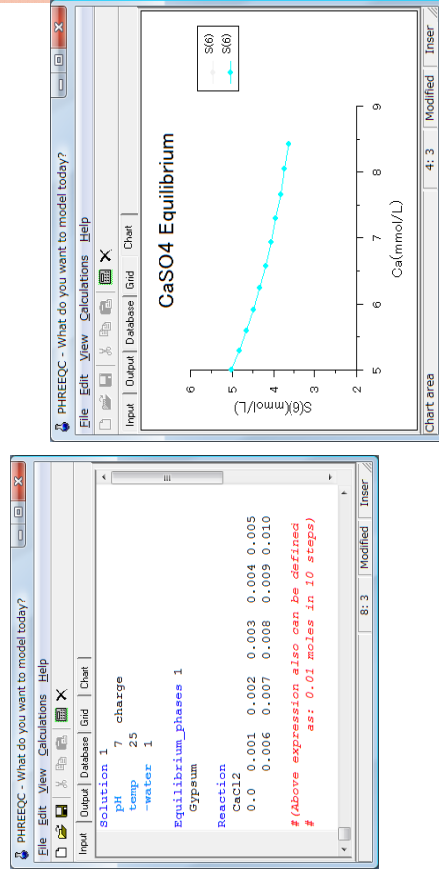
$$K = [\text{Ca}^{2+}]_{n-e} [\text{SO}_4^{2-}]_{n-e} = (0.00601 - \Delta x)(0.00501 - \Delta x) = 10^{-4.6} \\ \Rightarrow \Delta x = 0.00048 \text{ (mol/l)}$$

$$[\text{Ca}^{2+}]_{n-e} = 0.00601 - \Delta x = 0.00553 \text{ (mol/l)}$$

$$[\text{SO}_4^{2-}]_{n-e} = 0.00501 - \Delta x = 0.00453 \text{ (mol/l)}$$

Results

Q 3: Using the PHREEQC code.



PHREEQC: General description

PHREEQC can simulate:

(i) equilibrium reactions

- o aqueous complexation (reactions in pore-water)
- o cation exchange (on clay, organic matter, etc.)
- o surface complexation (e.g. iron oxyhydroxides)
- o mineral dissolution/precipitation, etc.

PHREEQC: General description

Example equilibrium reaction

$\text{Na}^+ - \text{Ca}^{2+}$ cation exchange process:



where X⁻ represents the amount of exchange sites [mol/kg soil]

and the *law of mass action* gives:

$$K_{\text{Na/Ca}} = \frac{[\text{Na-X}][1/2\text{Ca}^{2+}]}{[\text{Na}^+][1/2\text{Ca-X}_2]}$$

where $K_{\text{Na/Ca}}$ is the equilibrium constant and “[¹⁶]” represents concentration

PHREEQC: General description

PHREEQC can simulate:

(ii) kinetic reactions

- degradation reactions (e.g., nitrogen cycle: $\text{Org-N} \rightarrow \text{NH}_4^+ \rightarrow \text{NO}_3^- \rightarrow \text{N}_2$)
- oxidation-reduction reactions (e.g., pyrite (FeS_2) oxidation)
$$\text{FeS}_2 + 15/4 \text{O}_2 + 7/2 \text{H}_2\text{O} \rightarrow \text{Fe}(\text{OH})_3 + 2 \text{SO}_4^{2-} + 4 \text{H}^+$$
- radioactive decay (uranium-238 chain: $\text{U-238} \rightarrow \text{U-234} \rightarrow \text{Th-230} \rightarrow \text{Ra-226}$)

PHREEQC: General description

Example kinetic reaction

Nitrogen transformation process:



with the first-order rate degradation reactions:

$$\frac{d\text{Org-N}}{dt} = -\mu_{\text{Org-N}} \text{Org-N} \quad \frac{d\text{NH}_4^+}{dt} = -\mu_{\text{NH}_4^+} \text{NH}_4^+ + \mu_{\text{Org-N}} \text{Org-N} \quad \frac{d\text{NO}_3^-}{dt} = -\mu_{\text{NO}_3^-} \text{NO}_3^- + \mu_{\text{NH}_4^+} \text{NH}_4^+$$

where $\mu_{\text{Org-N}}$, $\mu_{\text{NH}_4^+}$ and $\mu_{\text{NO}_3^-}$ are the first-order rate constants.

- ❖ In PHREEQC, the rate constant μ could be function of the temperature.

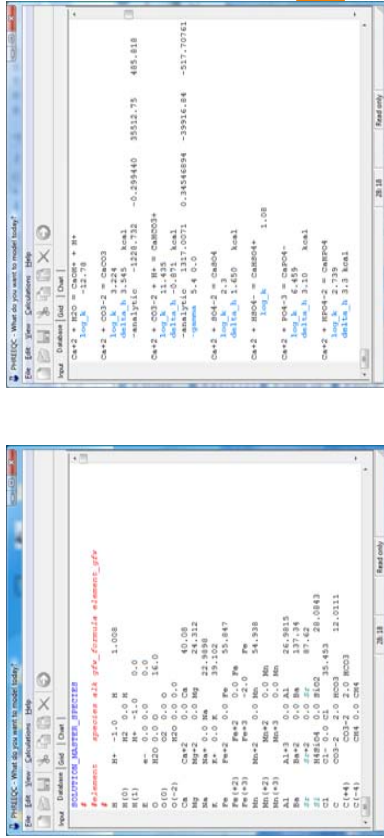
PHREEQC can simulate:

(iii) mixed equilibrium/kinetic reactions

- chain reactions (modeled as kinetic type) including adsorbed phase (modeled as equilibrium type), etc.

PHREEQC: General description

PHREEQC contains databases with chemical definitions (*species, complexes, mineral solubility, etc.*) and constants needed to do the calculations.



The convection–dispersion equation (CDE) in HP1

The general solute transport equation in HP1 is:

$$\frac{\partial \theta c_{i,l}}{\partial t} = \frac{\partial}{\partial x} \left(\theta D^w \frac{\partial c_{i,l}}{\partial x} \right) - q \frac{\partial c_{i,l}}{\partial x} - S c_{r,i} + R_i$$

where

- i ($\in 1, \dots, N_m$) is the aqueous species number (N_m is the total number of aqueous species),
- $c_{i,l}$: aqueous concentration phase [ML⁻³],
- θ : volumetric water content [L³L⁻³],
- D^w : diffusion-dispersion coefficient [L²T⁻¹],
- q : water flux [LT⁻¹],
- S : water uptake term [L³L⁻³T⁻¹],
- $c_{r,i}$: the concentration of the sink term [ML⁻³], and
- R_i : general source/sink term due to geochemical reactions [ML⁻³T⁻¹].

Simulation Tool



Simulating water flow, transport and biogeochemical reactions in environmental soil quality problems

HP1

A Coupled Numerical Code for
Variably Saturated Water Flow,
Solute Transport and
Biogeochemistry
in Soil Systems

HP1

Flow and Transport Model
HYDRUS-1D 4.0

Biogeochemical Model
PHREEQC-2.4

Source/Sink Term R_i

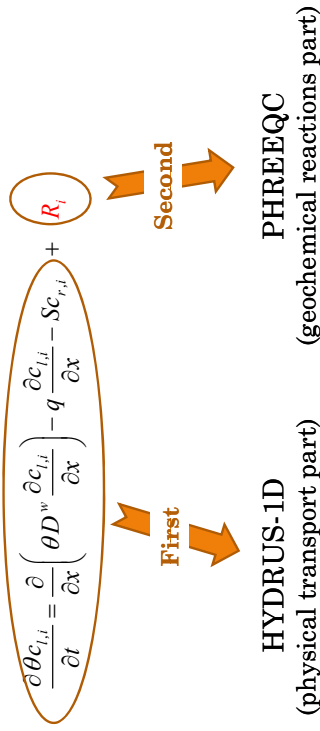
The general source/sink term R_i represents any change of the soil system due to the biogeochemical processes, and namely:

- equilibrium homogeneous and heterogeneous reactions,
- kinetic homogeneous and heterogeneous reactions,
- mixed equilibrium - kinetic reactions.

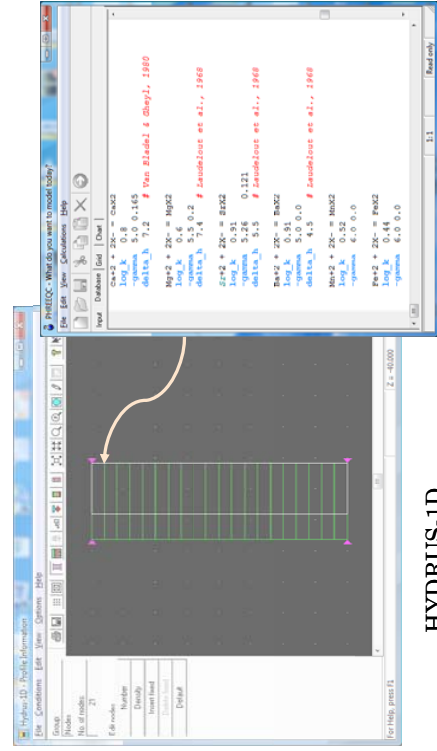
HP1 coupling procedure

- Coupling method: non-iterative sequential approach (weak coupling)
- Within a single time step:
 - Solve water flow equation (HYDRUS)
 - Solve heat transport equation (HYDRUS)
 - Solve CDE for solute transport like inert components (HYDRUS)
 - Calculate speciations, equilibrium reactions, kinetic reactions, etc. ... for each cell (PHREEQC)

HP1 coupling procedure



HP1 coupling procedure



HYDRUS-1D
(nonreactive solute transport in the profile)

At each node, PHREEQC evaluates chemical reactions.

HP1 coupling procedure

- The total mass balance for the geochemical reactions at each node is considered in the PHREEQC module
- Since explicit time scheme is used, small Δt is necessary
- The coupling procedure is inefficient and time-consuming

Application of the code

- Transport of heavy metals (Zn^{2+} , Pb^{2+} , and Cd^{2+}) subject to multiple cation exchange
- Transport with mineral dissolution of amorphous SiO_2 and gibbsite ($\text{Al}(\text{OH})_3$)
- Heavy metal transport in a soil with a pH-dependent cation exchange complex

Application of the code

- Infiltration of a hyperalkaline solution in a clay sample including kinetic precipitation-dissolution of kaolinite, illite, quartz, calcite, dolomite, gypsum, hydrotalcite, and sepiolite
- Long-term transient flow and transport of major cations (Na^+ , K^+ , Ca^{2+} , and Mg^{2+}) and heavy metals (Cd^{2+} , Zn^{2+} , and Pb^{2+}) in a soil profile
- Kinetic biodegradation of NTA /nitriiotriacetate/ (biomass, cobalt)

Application of the code

- Transport of TNT /trinitrotoluene; trotyl/ and its daughter products
- Uranium transport in agricultural field soil (example considers surface complexation reactions due to Fe-oxides)
- Horizontal infiltration and adsorption of piggery effluent

HP1 summary

- Tool that integrates physical – chemical – biological aspects of chemical migration in soils
- Multiple biogeochemical reactions
- Complex flow fields
- Spatial variability

HP1 summary

- Interpretation of laboratory and field data
- Tool for understanding qualitatively and quantitatively trends in data
- Sensitivity analyses
- Permits systematic evaluation of impact of initial and boundary conditions on model output

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